SFIT4 – a comprehensive tool to analyze atmospheric spectra measured by ground-based FTIR spectroscopy.

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Boulder, 2019



Absorption spectroscopy



Measurements in solar or lunar absorption and emission possible



Absorption spectroscopy



Measurements in solar or lunar absorption and emission possible

Received spectrum given by the sum of all absorption along the path of sight.



Envelope defined by band filter 3900 - 4400 cm^{-1}



Absorption spectroscopy



Measurements in solar or lunar absorption and emission possible

Received spectrum given by the sum of all absorption along the path of sight.



Microwindow containing a CO-line



Absorption spectroscopy



Measurements in solar or lunar absorption and emission possible

Received spectrum given by the sum of all absorption along the path of sight.



Contribution of gases and solar lines not calculated for this particular spectrum



Principle of a Fourier transform spectrometer



- Overview SFIT4
- Properties of the radiative transfer model
- Retrieval
- Error/Sensitivity calculation





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Overview of sfit4 package

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STRATOSPHERIC NO, AND H,O MILLING RATIO PROFILES FROM HIGH RESOLUTION INFRARED SOLAR SPECTRA

USING NONLINEAR LEAST SQUARES

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- sfit development started in the early 90ies
- aim: fast and small code to be less dependent on large computers
- first version of sfit2 ready in 1998

- SFIT4 started in 2010, first official version in 2014
- cooperation of many groups operating FTIR: NCAR Boulder, U Toronto, BIRA Brussels and U Bremen



Overview of SFIT4



Atmosphere and auxilliary data are representation of reality



FORWARD MODEL F:

- y spectrum
- x_g atmospheric state
- x_b auxiliary parameters
 - ϵ noise on spectrum

$$y = F(x_g, x_b) + \epsilon$$

RETRIEVAL: recipe to modify x_g and x_b

- METHODS:
 - optimal estimation
 - Tikhonov-Phillipsregularization
- CAVEAT: A PRIORI information necessary due to lack of information in spectrum

y – spectrum

$$x_g$$
 – atmospheric state

$$y = F(\underbrace{x}_{=(x_g,x_b)}) + \epsilon$$

 x_b – auxiliary parameters

- for atmosphere:
 - line-by-line model using the Voigt line shape
 - raytracing described by LBLATM
 - solar line parameters by Frank Hase
 - spectroscopic data from data bases, e.g. HITRAN
- for instruments:
 - line-shape effects: apodization function, phase shift, field-of-view
 - frequency shift of instrument
 - zero offsets due to non-linearity of detector



v – spectrum x_q – atmospheric state x_b – auxiliary parameters $y = F(x) + \epsilon$

 $=(x_a,x_b)$

statevector x contains everything which could(!) be retrieved

$$X = (\underbrace{X_{Z=1,...,n}^{\text{VMR gas 1}}, X_{Z=1,...,n}^{\text{VMR gas 2}}, \dots, \underbrace{\nu_{\text{F-AXIS SHIFT}}, \nu_{\text{SOLAR SHIFT}}, \dots}_{X_{b}, \text{aux. parameters}})$$

 x_a is normalized or logarithmic for numerical reasons, i.e.

$$x_g = x_{ ext{Atmosphere}} / x_{ ext{A}}$$
 or $x_g = \log(x_{ ext{Atmosphere}})$



Atmospheric model Radiance calculated by

> $I = B(\infty) \exp(-\tau(0,\infty)) + \int_0^\infty \underline{\alpha(z')B(z')} \exp(-\tau(0,z'))dz'$ Emission of laver z $\tau(0,z) = \int_0^z \alpha(z') dz' \qquad \alpha(z) = \sum_{l=1}^N x_{a,l}(z) \alpha_l(z)$



B(z) Planck function Emission by Kirchhoff's law

$$e(\nu) = \alpha(\nu, P, T)B(\nu, T)$$

Transmission $\in [0, 1]$

$$T(\mathbf{0}, z) = \exp(-\tau(\mathbf{0}, z))$$

 $I = B(\infty) \exp(-\tau(0,\infty))$

Atmospheric model Radiance calculated by

$$)) + \int_{0}^{\infty} \underbrace{\alpha(z')B(z')}_{\text{Emission of lawyr, z}} \exp(-\tau(0,z'))dz'$$



$$\tau(0,z) = \int_0^z \alpha(z') dz' \qquad \alpha(z) = \sum_{l=1}^N x_{a,l}(z) \alpha_l(z)$$



Atmospheric model Radiance calculated by

a(z)B(Z)

a(z)B(Z)

Instrument

modeled spectrum

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$$I = B(\infty) \exp(-\tau(0,\infty)) + \int_{0}^{\infty} \underbrace{\alpha(z')B(z')}_{\text{Emission of layer }z} \exp(-\tau(0,z'))dz'$$

$$\tau(0,z) = \int_{0}^{z} \alpha(z')dz' \qquad \alpha(z) = \sum_{l=1}^{N} x_{a,l}(z)\alpha_{l}(z)$$
top of atmosphere
$$a(z)B(z)$$

0.6

0.5∟ 860

865 870 875 885 890 895 900

880

Wovenumber [em=1] 8/ 22

The absorption cross section

 The frequency dependent absorption cross section can be written as

$$\alpha(\nu) = S(\nu_0, T)L(\nu_0, \nu, n, P, T)$$
(1)

with S the transition intensity or source strength and L the line shape normalized to 1.

- S is the intensity of the line observed
- L contains the effects of the environment upon the observed molecule.



The absorption cross section – The intensity S

Intensity for transistion from state *i* to state *j*, S_{ij} is a quantity dependent on the transition and temperature

$$S_{ij} = C rac{\mu_{ij}^2}{TQ(T)} \left(e^{-rac{E_j}{k_B T}} - e^{-rac{E_i}{k_B T}}
ight)$$

- HITRAN contains intensity at 296K, the energy of the lower and the higher state
- partition function Q(t) calculated using TIPS method (part of HITRAN).
- extrapolation of intensity to arbitrary temperatures

$$S(T) = S(T_0) rac{T_0 Q(T_0)}{T Q(T)} \exp \left[rac{E_i + E_j}{2K_B T_0} \left(1 - rac{T_0}{T}
ight)
ight]$$



The absorption cross section – the line shape L

Active molecules: Molecules causing emission or absorption Perturbers: or buffering gas or bath, molecules not absorbing

Translational effects

- caused by the (thermal) movement of the molecules
- Velocity distribution is Maxwell Boltzmann if not disturbed
- Line shape described by a Gauss function L_G

Voigt function

Collisional effects

- Caused by interactions of molecules
- Dephasing of radiation -> limiting correlation between different times
- line shape described by a Lorentz function L_L

 $L_V(\nu) = L_G \star L_L$



The absorption cross section – the line shape L



Line shape described by a Gauss function L_G

Voigt function

 $L_V(\nu) = L_G \star L_L$



The absorption cross section – the line shape L

 Line width of Lorentz function increases
 proportional to pressure.
 Consequence of uncertainty relation

$\Delta E \Delta t \leq \hbar$

 Width of Gaussian part is roughly constant in altitude but proportional to frequency

NOTE: Altitude information is only available as long as the Lorentz part dominates



The absorption cross section – the line shape L



Calculation of the K matrix

For the retrieval and sensitivity study we need the first derivative, the so-called Jacobian of the forward model, also called the weighting function matrix K

Remember: $x = (x_g, x_b)$

$$y = F(x)$$

$$K_{g,b} := \frac{\partial F}{\partial x_{g,b}}\Big|_{x_{g,b}^0}$$

Using the K matrix, the forward model can be linearized:

$$F(x) = F(x_0) + K(x_0)(x - x_0) + O\left((x - x_0)^2\right)$$

using $\tilde{y} = y - y_0$
 $\tilde{x} = x - x_0$
 $\rightsquigarrow \tilde{y} = K\tilde{x}$



Calculation of the K matrix

Two approaches:

1. Perturbation: for each row of the K matrix:

$$K_i = rac{F(x + \Delta x) - F(x)}{\Delta x}$$

-> order N²

 Semi analytic K-matrix calculation, exploits the fact that the transmission in of a layer only changes if the VMR underneath this layer changes. The spectrum and the K-matrix are calculated in two runs from TOA to the ground → order N



Inversion of RTM *F* is an ill-posed problem

- In the presence of noise, information content of spectra in nadir/zenith geometry is limited, not more than a few bit
- influence of noise grows exponentially
- but: for numerical reasons, atmospheric bins should be small
- \rightsquigarrow number of layers higher than information content \rightsquigarrow regularization is needed

In SFIT4 two methods are implemented

- Optimal estimation (in the version of Rodgers, 2000)
- Tikhonov-Phillips-Regularization (various authors)



Inversion using Bayes theorem

p(y|x) - conditional probability distribution

$$p(x|y)p(y) = p(y|x)p(x)$$
 Bayes theorem
 $\rightsquigarrow p(x|y) = \frac{p(y|x)p(x)}{p(y)}$

We choose all probabilities to be Gaussian Why?

- ▶ product of two Gaussian functions is Gaussian → p(x|y) is Gaussian
- mode = mean
 - $\rightsquigarrow \int p(y)$ does not need to be known



Inversion using Bayes theorem

$$p(y|x) = C_L \exp(-(y - F(x))^T S_{\epsilon}^{-1}(y - F(x)))$$

$$p(x) = C_A \exp(-(x - x_A)^T S_A^{-1}(x - x_A))$$

$$\rightsquigarrow p(x|y) = C \exp(-(x - \hat{x})^T S^{-1}(x - \hat{x}))$$

getting the mode (=mean) by finding the extrema of the argument:

$$\hat{x} = \arg\min(\underbrace{(y - F(x))^T S_{\epsilon}^{-1} (y - F(x)) + (x - x_A)^T S_A^{-1} (x - x_A)}_{\text{cost function}})$$

$$\hat{x} = x_A + (S_A^{-1} + K^T S_{\epsilon}^{-1} K)^{-1} [K^T S_{\epsilon}^{-1} (y - K x_A)]$$

$$S^{-1} = K^T S_{\epsilon}^{-1} K + S_A^{-1}$$



Inversion using Tikhonov-Phillips-Regularization by minimizing:

$$\hat{x} = \arg\min(\underbrace{||P(y - F(x))||}_{\text{data misfit}} + \lambda \underbrace{||R(x - x_A)||}_{\text{Regularization}})$$

 λ – regularization strength

Solution by

$$\hat{x} = x_A + (R^{-2} + K^T P^{-2} K)^{-1} [K^T P^{-2} (y - K x_A)]$$

 \rightsquigarrow equivalent to optimal estimation with $R^2 = S_A$ and $P^2 = S_\epsilon$



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 \sim equivalent to optimal estimation with $R^2 = S_A$ and $P^2 = S_\epsilon$ NOTE: Optimal Estimation and Tikhonov-Phillips-Regularization both have Gaussian statistics



Inversion of the RTM

non-linear forward models

Two iteration methods implemented:

- 1. Gauss-Newton iteration
 - weakly non-linear models
 - iteration only successful if started near the solution

 $x_{i+1} = x_{\mathcal{A}} + (S_{\mathcal{A}}^{-1} + K_i^T S_{\epsilon}^{-1} K_i)^{-1} K_i^T S_{\epsilon}^{-1} (y - F(x_i)) + K_i (x_i - x_{\mathcal{A}})$

- 2. Levenberg-Marquardt iteration
 - moderatly non-linear models
 - compromise of Gauss-Newton (quick but unstable) and steepest descend iteration (slow but stable)

 $x_{i+1} = x_{\mathcal{A}} + ((1+\gamma)S_{\mathcal{A}}^{-1} + K_i^T S_{\epsilon}^{-1} K_i)^{-1} K_i^T S_{\epsilon}^{-1} [(y - F(x_i)) + K_i(x_i - x_{\mathcal{A}})]$

 γ is the weighting between Gauss-Newton and steepest descent. γ decreases when iteration successful (cost function gets lower), else is increased

 $\leadsto \gamma$ start with a high value and is continually decreased



The averaging kernel matrix

One of the most important quantities beside the result is the sensitivity of the retrieval, or the avarging kernels A:

$$A := \frac{\partial \hat{x}}{\partial x} \\ = \frac{\partial \hat{x}}{\partial F} \frac{\partial F}{\partial x}$$

comparing with and identifying y = Kx:

$$\hat{x} = x_{A} + (S_{A}^{-1} + K^{T}S_{\epsilon}^{-1}K)^{-1}K^{T}S_{\epsilon}^{-1}(y - Kx_{A})$$
$$\hat{x} = x_{A} + \underbrace{(S_{A}^{-1} + K^{T}S_{\epsilon}^{-1}K)^{-1}K^{T}S_{\epsilon}^{-1}}_{=D}K(x - x_{A})$$

we find a linarisation of the retrieval using A = DK:

$$\hat{x} = x_A + A(x - x_A) = (I - A)x_A + Ax$$



- Can the measured spectrum be modeled?
- Is the result sensible?
- Correlation of retrieved quantities
- Error/Sensitivity of retrieved quantities



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artifacts in spectrum?

- origin?
- is the failure in modeling relevant?





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- artifacts in spectrum?
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- result different from a priori?
- structure of the retrieved result ok?
- error of the retrieval sensible?





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 the correlation of different entries of the state vector is given by the AVK matrix





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- Ozone measurements using a millimeterwave instrument
- Measurements independent of weather and light
- Altitude range about 20 -60 km altitude

- Can the measured spectrum be modeled?
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Change of Ozone VMR in %



Upper panel: measurements of Ozone VMR during a solar proton event

Lower panel: modelling of the same event using a chemical transport model and realistic ion fluxes

- Can the measured spectrum be modeled?
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Change of Ozone VMR in %



Upper panel: measurements of Ozone VMR during a solar proton event

Lower panel: the model has been reculatulated using the measurement model

What would the instrument see, if the model would be true?

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- Can the measured spectrum be modeled?
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$$\begin{aligned} \hat{x} - x &= (A - I)(x - x_A) \\ &+ G_y \epsilon \\ &+ G_y \Delta F(x, b, \hat{b}) \\ &+ G_y K_b(b - \hat{b}) \end{aligned}$$

Smoothing error Retrieval noise Forward model error model parameter error

compare Rodgers (2000)



$$\hat{x} - x = (A - I)(x - x_A) + G_y \epsilon + G_y \Delta F(x, b, \hat{b}) + G_y K_b(b - \hat{b})$$

Smoothing error Retrieval noise Forward model error model parameter error

smoothing error only accessible when covariance of real ensemble is known.

retrieval noise caused by noise on spectrum.

Error due to use of wrong forward model. Difficult to assess if true forward model is not known.



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$$\hat{x} - x = (A - I)(x - x_A) + \frac{G_y \epsilon}{G_y \Delta F(x, b, \hat{b})} + G_y \mathcal{K}_b(b - \hat{b})$$

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